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REE behavior and effect factors in AMD-type acidic groundwater at sulfide tailings pond, BS nickel mine, W.A.

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Abstract: AMD(Acid Mine Drainage)-type acidic groundwater (pH < 4) from oxidizing sulfide tailings in BS nickel mine (Western Australia) is of higher total rare earth element(REE) contents and Ce enrichment (PAAS normalization), different from setting groundwater (pH > 6.5, with lower total REE contents, Ce depletion). While the AMD contaminated groundwater (pH=4.0-6.5) around tailings pond is characterized by transition from acidic to setting groundwater in total REE content, and associated with Ce depletion (like setting groundwater). The light REE in all type groundwater shows up depletion, but its depleted extent in acidic groundwater is more remarkable. This work indicates that REE behavior in AMD-type acidic groundwater is controlled mainly by pH value and metal (Al, Mn and Fe) contents. And the critical pH value that affects REE behavior in ground acidic water would be 4, lower than the previous value (pH=5) that has been believed prevalently in surface acidic waters. The pH could affect REE behavior in groundwater by controlling the solubility of metal (Al, Mn and Fe) hydroxides and the valence of cerium. Finally, light REE depletion in acidic groundwater may due to element affinity. High content Al (affinity with heavy REE) and low content Fe (affinity with light REE) may lead to heavy REE enrichment while light REE relative depletion in water.

Key words: REE pattern; acid mine drainage; sulfide tailings; groundwater contamination

1 Introduction

Rare earth element(REE) can record subtle geochemical processes in natural systems, and has been extensively applied to earth science field. However, REE geochemical behavior in aquatic systems is a relatively new field of study. In recent years, many studies have focused on REE concentration, differentiation and effect factors in surface acidic waters, e.g. river, lake and mine surface water[1–5], but few papers published have been involved in ground acidic waters[6–7], especially AMD(Acid mine drainage)-type acidic groundwater.

There are a volume of solid wastes (including mill tailings, waste rock, etc.) produced during the period of metal sulfide mining activity. When the sulfide minerals remaining in the solid wastes expose to earth surface, they react with H_2O and O_2 in atmosphere, resulting in the generation of sulfuric acid. The acid drainage/AMD

will happen only if the neutral/buffer capacities of the solid waste itself and environment were overwhelmed by its acid generation potentials. In contrast with the surface acidic water rich in oxygen, the ground acidic water which was formed by acid drainage infiltrating downward may be almost under anaerobic condition, meaning that biological and biochemical action may be of no consideration relatively in groundwater. Hence, there should be a difference in REE distribution patterns between surface and ground acidic water. In this study, REE behavior and effect factors in AMD-type acidic groundwater are studied through sampling and testing groundwater from drilling holes, which are placed nearby and beyond tailing bond in BS nickel sulfide mine, Western Australia.

2 Setting

BS nickel sulfide mine is near to Kalgoorlie, Western

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Australia. In BS mine, milling wastes, massive sulfide tailings (approximately 58 000 t/a) and disseminated sulfide tailings (approximately 325 000 t/a) are stored in MST and DST pond, respectively (Fig.1). Acidified groundwater has been recognized around the tailings pond three years after its establishment. A recent assessment of the tailings disclosed that the massive sulfide tailings were totally classified as potentially acid forming, and approximately 16% by volume of the disseminated sulfide tailings held a potential to be acid generating[8].



Fig.1 Location map showing sampling sites: MST—Massive sulfide tailings pond; DST—Disseminated sulfide tailings pond; EP—Evaporation pond; T_{1-12} —Site and number of monitoring bores; B_{2-5} —Site and number of active pump bores

In the vicinity of tailings pond, there is a clayey ferricrete cover (4–20 m deep), which is underlain by saprolitic clay with minor ferruginous bands (62–79 m deep). Weathered depth (approached to Archaean mafic and ultramafic rock base) is up to 62–126 m. Surface runoff is to Southwest then South into a poorly defined channel outside the mine site. Groundwater is also interpreted to flow to South under low hydraulic gradients, and natural water table was 25–29 m below ground level.

3 Sampling and test results

3.1 Sampling

Groundwater samples were collected from monitoring bores (T_{1-12}) nearby the pond and active pump bores (B_{2-5}) beyond the pond (Fig.1). Groundwater pH (and electrical conductivity, static water level and temperature) was measured on site. A portion of each water sample (filtered through 0.45 μ m nitrocellulose filter, then acidified with ultra pure nitric acid) was analyzed for REE and trace elements by ICP-MS and ICP-AES in department of Applied Chemistry, Curtin University of Technology (Western Australia). REE test values of groundwater samples were normalized using Post-Archaean Australian Shale (PAAS), and Cerium anomaly (Ce/Ce^{*}) was calculated using the formula:

 $A(Ce/Ce^*)=N(Ce)/[N(La)\times N(Pr)]^{0.5}$

where N denotes PAAS-normalized value.

3.2 Results

Based on pH values, the groundwater samples can be divided into three types (Table 1): type , pH < 4, acidic groundwater (including sample T₂, T₃, T₅, T₇, T₁₀), collected in the vicinity of pond; type , pH=4.0—6.5, acid contaminated groundwater (including T₄, T₆, T₈, T₁₂), also collected in the vicinity of pond; type , pH > 6.5, setting groundwater, no acid contamination. Type samples include B₂, B₃, B₄, B₅ (sampling sites away from the tailing pond), and T₁, T₉, T₁₁ (sampling sites near to the pond, but their water samples are of pH > 6.5, hence these samples are classified as type).

Three-type groundwaters display different characteristics in solved REE distribution patterns (PAAS normalization) (Fig.2 and Table 1): 1) REE concentration in type $((75.5-671) \times 10^{-9})$, average 301.3×10^{-9}) is of 1–2 orders of magnitude higher than that in type $((4.44-32.8) \times 10^{-9}, \text{ average } 9.51 \times 10^{-9});$ while REE concentration in type is of greater variation ((6.26–677) $\times 10^{-9}$, average 202.49 $\times 10^{-9}$), and transition from type to type . 2) cerium anomaly in has a vaster variation (Ce/Ce^{*}=0.655-1.66, type average 1.12), but most of the samples (except T_2) show weak-middle Ce positive anomaly; while types and

Grossly speaking, the acidic groundwater (type) in the mine is characterized by REE high concentration, Ce enrichment, and light REE depletion. In addition, it has a weak-middle REE enrichment (Fig.2).

4 Discussion

4.1 REE contents

Many previous studies indicated that REE concentration is higher in acidic water than in normal water, and that pH plays an important role in controlling the concentration of solved REE in water, and its critical

Sample No.	pH	w(Al)/ 10 ⁻⁶	w(Mn)/ 10 ⁻⁶	$w(Fe)/10^{-6}$	w(La)/ 10 ⁻⁹	w(Ce)/ 10 ⁻⁹	w(Pr)/ 10 ⁻⁹	w(Nd)/ 10 ⁻⁹	w(Sm)/ 10 ⁻⁹	w(Eu)/ 10 ⁻⁹	w(Gd)/ 10 ⁻⁹
T ₁₀	3	147.0	3.23	< 0.2	78.7	327	38.4	176	45.7	16.1	54.8
T_2	3.3	49.9	66.2	< 0.2	373	606	122	474	119	49.3	145
T ₃	3.3	114.0	13.6	< 0.2	96	499	50	234	65.3	22.5	74.7
T ₇	3.19	72.6	2.03	0.289	50.8	129	18.1	64.9	14	5.64	16.8
T ₅	3.91	10.1	11.1	0.398	60.2	178	30.7	132	33.5	11.4	35.9
T ₁₂	4.8	1.69	14.3	0.294	416	288	121	498	123	47.3	159
T_8	5.28	0.857	1.59	< 0.2	61.8	33.1	14.3	66.4	14.8	6.19	21.6
T_6	6.08	0.4	0.165	0.302	7.03	7.65	2.45	8.85	2.19	0.292	2.1
T_4	6.48	0.418	2.08	< 0.2	34.2	28	8.07	35.9	6.28	2.25	7.82
T_1	6.7	0.51	7.58	< 0.2	24.4	67.5	8.37	31	6.81	2.33	9.05
Т9	7.41	0.393	0.09	< 0.2	5.98	6.33	1.97	8.34	1.48	0.919	1.4
T ₁₁	6.96	0.262	0.042	< 0.2	5.69	4.82	1.91	8.9	1.21	0.325	1.26
B_2	6.6	0.242	0.038	0.214	4.79	6.02	1.71	9.45	1.34	0.325	1.26
B ₃	6.51	0.299	0.572	0.252	7.18	7.62	2.73	11.1	2.24	0.91	1.96
B_4	6.75	0.36	0.249	0.511	4.88	6.24	2.13	9.07	1.67	0.182	1.46
B_5	6.96	0.308	0.119	0.479	5.25	5.32	1.86	7.06	1.79	0.107	1.26
Sample No.	w(Tb)/ 10 ⁻⁹	w(Dy)/ 10 ⁻⁹	w(Ho)/ 10 ⁻⁹	w(Er)/ 10 ⁻⁹	w(Tm)/ 10 ⁻⁹	w(Yb)/ 10 ⁻⁹	w(Lu)/ 10 ⁻⁹	w(<u>SREE</u>)/ 10 ⁻⁹	Cerium anomaly	N(Gd/La)	N(La/Lu)
T ₁₀	9.88	5.6	12.3	32.2	4.45	26.7	3.45	232	1.37	5.71	0.314
T_2	31.6	172	37.3	94	13.7	73.6	9.57	671	0.655	3.19	0.283
T_3	14.6	93.6	20.9	54.5	7.57	43.8	5.89	359	1.66	6.38	0.315
T_7	3.080	18.1	2.82	10.9	1.36	8.93	0.929	75.5	0.982	2.71	0.277
T ₅	8.25	45.1	9.05	24.6	3.17	18	2.42	169	0.955	4.89	0.309
T ₁₂	29.5	172	37	97.3	13	70	9.47	677	0.296	3.13	0.295
T_8	4.4	22.9	4.93	16.2	1.79	9.66	1.6	94.9	0.257	2.87	0.362
T_6	0.424	1.24	0.284	0.431	0.145	0.408	0.118	6.26	0.425	2.45	0.33
T_4	1.83										
T_1	1.05	8.04	1.49	3.59	0.532	2.63	0.396	31.8	0.389	1.87	0.301
	1.38	8.04 7.43	1.49 1.47	3.59 3.59	0.532	2.63 3.29	0.396	31.8 32.8	0.389	1.87 3.04	0.301
Т9	1.38 0.406	8.04 7.43 1.11	1.49 1.47 0.133	3.59 3.59 0.487	0.532 0.514 0.128	2.63 3.29 0.818	0.396 0.436 0.079	31.8 32.8 6.19	0.389 1.09 0.426	1.87 3.04 1.92	0.301 0.344 0.25
Т9 Т ₁₁	1.38 0.406 0.248	8.04 7.43 1.11 0.888	1.49 1.47 0.133 0.112	3.59 3.59 0.487 0.694	0.532 0.514 0.128 0.144	2.63 3.29 0.818 0.486	0.396 0.436 0.079 0.077	31.8 32.8 6.19 4.92	0.389 1.09 0.426 0.337	1.87 3.04 1.92 1.82	0.301 0.344 0.25 0.217
T ₉ T ₁₁ B ₂	1.38 0.406 0.248 0.287	8.04 7.43 1.11 0.888 0.82	1.491.470.1330.1120.225	3.59 3.59 0.487 0.694 0.266	0.532 0.514 0.128 0.144 0.144	2.63 3.29 0.818 0.486 0.486	0.396 0.436 0.079 0.077 0.175	31.8 32.8 6.19 4.92 4.66	0.389 1.09 0.426 0.337 0.485	1.87 3.04 1.92 1.82 2.16	0.301 0.344 0.25 0.217 0.492
T_9 T_{11} B_2 B_3	1.33 1.38 0.406 0.248 0.287 0.21	8.04 7.43 1.11 0.888 0.82 1.85	1.49 1.47 0.133 0.112 0.225 0.301	3.59 3.59 0.487 0.694 0.266 0.804	0.532 0.514 0.128 0.144 0.144 0.163	2.63 3.29 0.818 0.486 0.486 0.487	0.396 0.436 0.079 0.077 0.175 0.215	31.8 32.8 6.19 4.92 4.66 7.82	0.389 1.09 0.426 0.337 0.485 0.397	1.87 3.04 1.92 1.82 2.16 2.24	0.301 0.344 0.25 0.217 0.492 0.534
T9 T11 B2 B3 B4	1.38 1.38 0.406 0.248 0.287 0.21 0.289	8.04 7.43 1.11 0.888 0.82 1.85 1.31	1.49 1.47 0.133 0.112 0.225 0.301 0.245	3.59 3.59 0.487 0.694 0.266 0.804 0.808	0.532 0.514 0.128 0.144 0.144 0.163 0.091	2.63 3.29 0.818 0.486 0.486 0.487 0.489	0.396 0.436 0.079 0.077 0.175 0.215 0.118	31.8 32.8 6.19 4.92 4.66 7.82 5.71	0.389 1.09 0.426 0.337 0.485 0.397 0.447	1.87 3.04 1.92 1.82 2.16 2.24 2.45	0.301 0.344 0.25 0.217 0.492 0.534 0.525

Table 1 Analysis results of groundwater in drill holes, BS mine site*

 T_{1-12} —Drilling holes around tailing pond and sample number; B_{2-5} —Drilling holes away from pond and sample number.

value is pH=5. Below a pH of 5 in acidic water, REE concentrates in water facies. The reason is that REE behaves conservatively in acidic water, and the elements concentrate in water facies as liberal ion or complex with $SO_4^{2^-}$ (F⁻). While at pH > 5, REE tends to deposit though completing with $CO_3^{2^+}$ or $(CO_2)^-$, and be adsorbed by the metal (e.g. Al, Fe, Mn) hydroxide colloid, leading REE contents in waters to reduce[9–17]. And there is a negative correlation between REE content and pH

value[18-19].

This paper got similar results to the previous studies. As pH value is below 5, especially below 4, REE contents in groundwater increase greatly (Fig.3(a)). And Fig.4 displays a negative correlation between REE and pH in the groundwater (correlation coefficient r=-0.82), and a positive correlation between Al (Mn) and pH (r=0.92 and 0.87, respectively), with the exception of Fe (there is not remarkable correlation between Fe and pH).



Fig.2 PAAS normalization REE patterns in groundwater, BS mine: (a) pH < 4; (b) pH=4.0-6.5; (c) pH > 6.5

This reveals that REE contents in groundwater are controlled not only by pH, but also by Al, Mn contents. pH may control REE concentration through regulating adsorption-desorption of colloids (Al, Mn-hydroxides), meaning that REE adsorbed on the colloids in water can be released into water with pH decreasing (which leads colloid electronegativity to decrease), and vice versa[16]. These imply that pH is a dominant factor in controlling REE contents in groundwater.

In addition, the previous studies also revealed that REE contents in AMD-type acidic water, e.g. Odiel acid river (pH=3.25, REE 42.62 × 10^{-9}), Osamu Utsmi mine groundwater (pH=4.43, REE 1.046 × 10^{-6}), Metalliferous

Hills Cu-Pb-Zn mine stream (pH=3.1, REE 929 × 10^{-9}), and Chinese Sitai coal drainage (pH=3.52–3.75, REE (54.37–68.80) × 10^{-9})) have 2–3 orders of magnitude higher than those in normal water[1–2, 5–6]. In BS mine, the acidic groundwater (type, pH=3–3.91) should be classified as AMD-type acidic water, because its REE contents ((75.5–671) × 10^{-9}) are in the REE range ((42.62–1 046) × 10^{-9}) above AMD-type water and much higher than that in the normal setting groundwater.

4.2 Ce differentiation

pH value is one of the factors controlling Ce behavior in water, and the critical value is pH=5. At pH



Fig.4 Plots of correlation coefficient (r) of w(REE) with pH, w(Al), w(Mn) and w(Fe)

> 5, Ce is oxidized and changes from solved Ce^{3+} to dissoluble Ce^{4+} , then deposits as CeO_2 from water (BROOKINGS had ever certified the capability of CeO_2 sediment in nature waters)[20], resulting in a negative Ce anomaly in waters[21–23]. While in the pH < 5 acid water, Ce negative anomaly disappears and even it may show a small positive anomaly instead[11, 21].

Moreover, Ce behavior in water may relate to redox condition. LEYBOURNE et al[7] discovered that when they studied the difference of solved REE contents between surface and ground water in a mine site in Canada, Ce negative anomaly can be lowered down to 0.08 (NASC normalization) in surface water, while this Ce anomaly was not demonstrated in groundwater mostly. It may be due to the fact that, under surface aerobic condition, Ce is easier to be oxidized and deposited as CeO₂, causing Ce depletion in surface water; however, Ce in anoxic groundwater occurs as reduced Ce³⁺ instead and concentrates in water.

In BS mine site, Ce anomaly in groundwater is

controlled by pH, but the critical value is pH=4, relatively lower than the previous critical value (pH=5). In pH < 4 acidic groundwater, Ce shows weak-middle positive anomaly. In pH > 4, the positive anomaly tends to disappear(Fig.3(b)).

4.3 REE differentiation

Light REE is more active than heavy REE in water [22, 24], causing REE differentiation, but pH effect is more significant. Under high pH alkaline condition, light REE tends to be divorced from waters and concentrates into secondary minerals[13, 15]. And with pH increasing, the sequence of REE adsorption to particles (in lake) is: light REE > middle REE > heavy REE; contrary to pH decreasing, REE release has similar sequence[9]. Therefore, it is not surprising that the REE model with light REE depletion and heavy REE enrichment often appears in the alkaline water[10, 25], while light REE concentrates in the acid water relatively[4, 13].

In this study, all groundwater samples demonstrate light REE depletion (Fig.2), and in the pH < 4 acidic groundwater, the extent of light REE depletion is vaster ((Gd/La)_N values are higher in acidic water than in setting water, see Fig.3(c)). The unusual phenomenon of light REE depletion in acidic water may be due to the element affinity. MARMOLEJO-RODRIGUEZ et al[17] indicated that light REE tends to affiliate with Fe, while heavy REE prefers to do with Al in land water. In BS mine site, Al (and Mn) high content and Fe low content (Fe contents in many samples even are lower than test limitation, Table 1) in acidic groundwater may lead to heavy REE enrichment while light REE depletion relatively. Besides, WORRALL and PEARSON[26] had another explanation for light REE depletion in acidic water. It is due to the results of sulfide (e.g. pyrite) oxidizing, without involving in latter water-rock reaction. Therefore, the mechanism of light REE depletion in acidic water may be more complex, and it is necessary to do further study to get more persuasive explanation.

Besides, the middle REE enrichment, which exists in the general acidic water[1, 5], is also shown up in the AMD-type acidic groundwater in BS mine site, although its enrichment extent is markedly small (Fig.2).

5 Conclusions

1) REE behavior in AMD-type acidic groundwater is controlled by pH and metal (Al, Mn and Fe) contents.

2) The critical pH value that affects REE behavior in AMD-type acidic groundwater would be 4, lower than the critical value (pH=5) in general surface acidic water. The pH could affect REE behavior in groundwater by controlling the solubility of metal (Al, Mn and Fe) hydroxides and the valence of cerium. In pH < 4 acidic groundwater, metal hydroxides (with higher solubility), REE and Ce tend to concentrate in water, resulting in REE and Ce enrichment. However, in higher pH groundwater, the metal hydroxides deposit as colloid coagulation because of the solubility decreasing, leading to REE content decreasing. And Ce is easier to be oxidized and deposited as CeO_2 (Ce separates from other lanthanide by this way) under this condition, causing Ce depletion in water.

3) Light REE depletion can appear in AMD-type acid groundwater. One of the reasons may be due to the element affinity. High content Al (affiliated with heavy REE) and low content Fe (affiliated with light REE), like the studied acidic groundwater in BS mine, may cause heavy REE enrichment and light REE depletion relatively.

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